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RAPID VARIATIONS IN EARTH'S SATELLITE DRAG
AND PROPERTIES OF THE UPPER ATMOSPHERE

by

D. I. Irgashev
V. E. Chertoprud

(USSR)

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V. E. Chertoprud

SUMMARY

This note describes the correlation analysis of rapid variations of atmosphere density in the 165 - 215 km altitude range, determined by the fluctuations of the variation rates of anomalous periods of 11 American satellites tracked by radar.

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The satellites were observed in the period of solar activity minimum, in 1963 and 1964. The precision of density determination is very high; the interval of data smoothing constituted a few hours, the table spacing constituted 0.1 day, the total number of density estimates being ~360. It was shown in [1] that strong fluctuations in the Earth's atmosphere density take place even in periods of relatively moderate geomagnetic activity, and certain statistical estimates are brought out; however, statistical methods were not broadly applied.

An auto and cross-correlational method of analysis is presented in the current work. The values x of density, determined by satellite deceleration served as the initial material, alongside with the 3-hour geomagnetic index $a_p - y$, obtained from [2, 3]. The computed values of the autocorrelational and cross-correlational functions $r_r(xx)$, $r_r(yy)$, $r_r(xy)$ are represented by crosses in Figs 1 - 3. The theoretical auto-correlational and cross-correlational functions $\rho_r(xx)$, $\rho_r(yy)$, $\rho_r(xy)$ are also represented in the same figures for the simplest model:

$$\frac{d}{dt} x(t) + \lambda [x(t) - \langle x \rangle] = f(t), \quad \langle f \rangle = 0, \quad (1)$$

$$r_r(ff) = \exp [-\lambda |r|], \quad (2)$$

$$y(t) = a f(t - \Lambda) + \phi(t), \quad \frac{\langle y^2 \rangle}{a^2 \langle f^2 \rangle} = \epsilon. \quad (3)$$

The empirical and theoretical correlational functions agree for the values:

$$\Lambda = 16 \text{ days}^{-1}, \quad \lambda = 8 \text{ days}^{-1}, \quad \varepsilon = 2, \quad \Delta = 0.15 \text{ day} \quad (4)$$

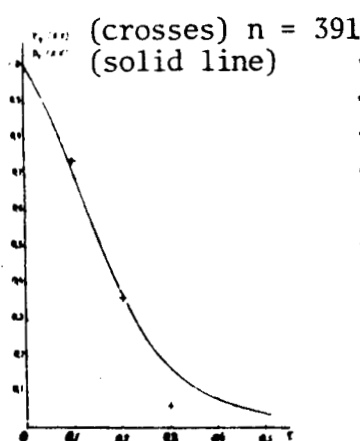


Fig. 1

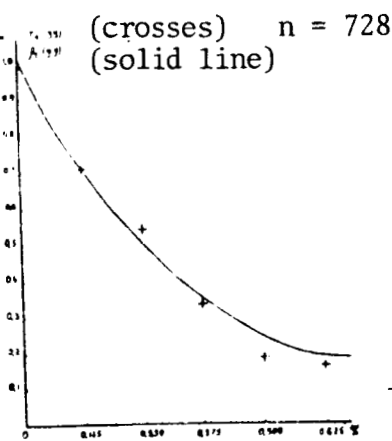


Fig. 2

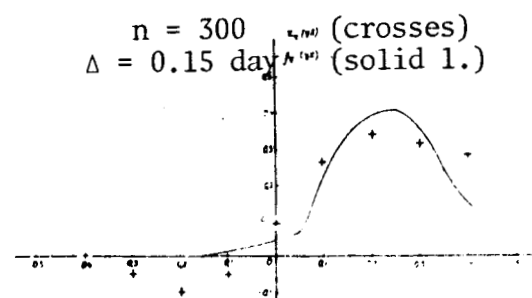


Fig. 3

Eq.(1) is the heat transfer equation of the thin layer, f is proportional to the variable heat flow, λ is proportional to the temperature conductance of the layer [4, 5] at the altitude where energy absorption takes place, linked with corpuscular streams. Eq.(2) is the simplest equation describing the link between the variable heat flow and the planetary geomagnetic index, Δ is the lag time on account of finite heat conductivity ($\Delta = 1/\lambda$).

Utilizing the standard atmosphere model [6] and the estimate (4), we obtain the value of the effective absorption height of the variable heat flow linked with geomagnetic disturbances:

$$z = 170 - 210 \text{ km.}$$

This effective height is substantially greater than the height of ultraviolet's absorption, which is $z = 110 - 120 \text{ km}$ [4, 5] and the actions upon the upper atmosphere of the variable ultraviolet and corpuscular radiations are divided in space by altitude.

These results have a preliminary character; they underscore the necessity of detailed study of rapid (hourly) variations of upper atmosphere density.

*** THE END ***

Shternberg (State) Astronomical
Institute in Moscow
June 1967

Translated by ANDRE L. BRICHANT
VOLT TECHNICAL CORPORATION, D.C.
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1. R. L. JACOBS. LMSC Tracking Note No.69, 1965
2. BARTELS..... [the rest of references is missing, the reprint being incomplete]